Nonlocal granular rheology: Role of pressure and anisotropy

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received 20 September 2013; accepted in final form 7 January 2014
published online 10 February 2014

PACS 47.57.Gc – Complex fluids and colloidal systems: Granular flow
PACS 83.80.Fg – Rheology: Granular solids
PACS 45.70.Mg – Granular flow: mixing, segregation and stratification

Abstract – We probe the secondary rheology of granular media, by imposing a main flow and
immersing a vane-shaped probe into the slowly flowing granulate. The secondary rheology is then
the relation between the exerted torque \( T \) and rotation rate \( \omega \) of our probe. In the absence of any
main flow, the probe experiences a clear yield-stress, whereas for any finite flow rate, the yield
stress disappears and the secondary rheology takes on the form of a double-exponential relation
between \( \omega \) and \( T \). This secondary rheology does not only depend on the magnitude of \( T \), but
it is also anisotropic —which we show by varying the relative orientation of the probe and main
flow. By studying the depth dependence of the three characteristic torques that characterize
the secondary rheology, we show that for counterflow, the dominant contribution is frictional-like
—i.e., \( T \) and pressure are proportional for given \( \omega \)— whereas for coflow, the situation is more
complex. Our experiments thus reveal the crucial role of anisotropy for the rheology of granular
media.

We still lack a full description of slow, dense flows of
granular media —given boundary conditions, what deter-
mines the flow rate? Granular flows differ in two impor-
tant ways from Newtonian flows. First, friction plays a
central role —when grains have persistent contacts, as
is the case in slow granular flows, inter-particle friction
provides the main channel by which energy can be dissi-
pated. Experiments in which the resistance to granular
flow is probed by controlling the stress on a single moving
boundary, such as in Couette, split-bottom or vane geo-
metries [1–3], find that the shear stresses are proportional
to the confining pressure, and that the ratio of shear to com-
pressive stress, which can be seen as an effective friction
coefficient, does not vary strongly with rate [1–5]. The
principal consequence from this rate independence is that
the stress plateaus for strain rates going to zero. In other
words, the stresses in slow granular flows are rate indepen-
dent, and therefore the stress is not sufficient to set the
strain rate.

The second ingredient is nonlocality. Several recent ex-
periments [6–8] and theoretical works [9–12] indicate that
for matter with granularity, the “fluidity” in location A,
i.e., the local relation between stress and strain rate, can
be strongly influenced by the flow in location B. Such
nonlocal behavior was first observed in the flow of emul-
sions [8], has also been observed in foams [13], and has
been modeled by a diffusive model for the fluidity, leading
to the introduction of a length scale which characterizes
the nonlocality [8,10,12]. Kamrin and coworkers have
adapted these ideas for frictional rheologies, capturing
granular Couette flows [9], and more recently, the full flow
profiles and height dependence of split bottom granular
flows [11,14–18].

In all these studies, there is a single source driving the
flow, and the nonlocality manifests itself via the spatial
flow profiles. However, as flow in location B influences
the fluidity in A, it is natural to study situations where
the driving of the main flow and the probing of the rhe-
ology are independent. Two examples of such granular
experiments which probe the “secondary rheology” con-
cern the sinking of a passive probe into a granular bed
that is stirred far away from the probe (in a split-bottom
cell) [7], and the rheology probed by rods submerged in and
dragged through a granular medium that is itself stirred
in a Couette geometry [6]. Recently, numerical studies
in which a plate is dragged through a 2D granular simple
shear flow have been performed [10]. In all cases, the yield stress experienced by the probes was found to vanish as soon as there is any external flow imposed, and the rate of the probe was proportional to the stirring rate, supporting the qualitative picture that flow in one location leads to mechanical agitations that govern the rheology in locations further away. Open questions concern: 1) the nature of the local rheology under stirring — in [7] the relation between stress and strain rate appears linear, whereas in [6,10] it appears exponential; 2) the effect of pressure — are shear stresses still proportional to pressure?; and 3) the role of anisotropy — are granular flows lead to anisotropy of the granular fabric [19,20] which should influence the fluidity of the material.

Here we probe these questions by characterizing the secondary rheology in a split bottom cell where the main flow is driven independently (fig. 1). Our experiments allow us to present full secondary rheological curves for a range of driving conditions and pressures. Moreover, we can probe the rheology when the probe moves with the main flow (coflow), against the main flow (counterflow) and also for probe rotation perpendicular to the main flow (fig. 1(b), (c)).

We find that, consistently with earlier work, whenever the main flow is present, the secondary rheology has a vanishing yield threshold [6,7,10]. The local rheology is anisotropic, and there are significant qualitative differences between coflow, counterflow, and perpendicular flow. Finally, by probing the flow at different depths, we find that the characteristic shear stresses are in some, but not all cases, proportional to the compressive stresses. Our results show that both local pressure and local anisotropy are crucial to describe the fluidity of granular media, thus stressing that these aspects need to be incorporated in future theories of slow granular flows.

Setup and protocol. — We create a smooth granular flow in a split-bottom shear cell [7,14–18] as shown in fig. 1(a). This cell consists of a cylindrical container of radius 80 mm, the bottom of which is split into a rotating disk of radius \( R_b = 60 \text{ mm} \) and a stationary outer ring. The cell is made of aluminum and the bottom and side walls are made rough by drilling dimples in them so as to obtain granular no-slip boundary conditions. The disk is driven by a micro-stepping motor at angular velocities \( \Omega \) ranging from \( 10^{-3} \) to 1 rps. The split-bottom cell is filled with a controlled mass of \( 1 \text{ mm} \) diameter glass beads up to the filling heights \( H \). We focus on large filling heights \( (H = 60 \text{ mm}) \), as the flow then predominantly takes place in a dome-shaped shear band buried in the granular medium [15–17,21] —this is the same flow geometry as used in [7].

To measure the rheological properties of this fluidized medium, we immerse a vane-shaped intruder into the material and couple it to an Anton Paar DSR301 rheometer. As illustrated in fig. 1(b), (c), the vane is made of crossed rectangular metallic plates of thickness 1 mm and of radius \( r \) by height \( h = 5 \times 5, 10 \times 10 \) or \( 20 \times 20 \text{ mm}^2 \). In most experiments, this probe is directly attached to the rheometer via a smooth and rigid shaft (fig. 1(b)), and we have checked that the residual torques picked up by the shaft are negligible in comparison to the signal from the vane. We have also carried out some experiments where we rotate the orientation of the vane as shown in fig. 1(c), and where we use a flexible tube with sufficiently large torsional stiffness to couple the vertical rheometer shaft to the horizontal vane shaft. The vertical distance between the center of the probe and the free surface of the material, \( z_m \) can be adjusted.

The rheometers native operation is in torque controlled mode, where torques ranging from \( -20 \) to \( 20 \text{ mNm} \) with a precision better than \( 1 \mu \text{mNm} \) are applied to our vane and the resulting deflection angle of the probe \( \phi(t) \) is measured. The precision on \( \phi \) is estimated to be significantly better than \( 10^{-4} \text{ rad} \), and in all experiments we deduce the average rotation rate \( \omega := \partial_t \phi(t) \) over deflection angles of \( 0.6 \text{ rad} \) or more. We define negative (positive) values of \( \omega \) to correspond to counter-rotation (co-rotation) with the bottom disk. Transients are short and all data presented are steady state values.

Prior to each measurement, the beads are vigorously stirred with a rod, after which the probe is immersed to its position \( z_m \). The system is then pre-sheared with \( \Omega = 0.5 \text{ s}^{-1} \) for 10 minutes, while the humidity of the system is lowered to values of \( 10–20\% \) relative humidity by closing the system with a plastic cover and injecting a dry air flow. This ensures a very good reproducibility of our rheological measurements.

Precession and yielding. — Ultimately, our goal is to probe and understand the relation between the probe rotation \( \omega \), disk rate \( \Omega \) and driving torque \( T \), \( \omega(\Omega, T) \), for a range of depths \( z_m \) and dimensions of the probe.
Fig. 2: (Color online) (a) The residual rotation rate $\omega_p$ of the probe in the absence of any torque on the probe is plotted as a function of depth $H - z$—curves are stretched exponential fits given by eq. (1). (b) Rotation rate $\omega$ of a $20 \times 20 \text{ mm}^2$ vane in the absence of a base flow ($\Omega = 0$) as a function of driving torque indicates a yield torque of $11.7 \pm 0.3 \text{ mNm}$.

Before delving into the full rheology, we first focus on the two single-forcing regimes: first the passive rotation of the probe due to residual flow in the system, $\omega(\Omega, T = 0)$, and second the rheology of the sand in the absence of stirring, $\omega(\Omega = 0, T)$.

We show in fig. 2(a) the behavior in the absence of a driving torque. As is well known, the rotation of the driving disk sets up a residual torsional flow in the medium $[7,16,17]$, that we characterize here by immersing one of our probes in the medium at depth $z_m$, setting the torque equal to zero, and measuring the ensuing precession of our probe, $\omega_p(z)$. For the three filling heights investigated, we find that $\omega_p(z)$ decreases with $z$ as a stretched exponential:

$$\omega_p(z)/\Omega = Ae^{-\left(\frac{H - z}{\xi}\right)^\alpha} + B,$$

where $A$ and $B$ are determined by the precession rates at $z = 0$ ($\omega_p$) and at $H - z = 0$ [$7$] $-\alpha \approx 1.5$ and $\xi$ represents a characteristic size of the main flow and is of the order of ten grain diameters. Moreover, the precession velocity is found to be independent of the probe size (not shown), and proportional to $\Omega$. From the fits of the precession velocity, we can get an estimate of the local strain rate of the flow: $\dot{\gamma}(z) \approx \partial \omega / \partial z$. For the filling heights used here, the shear band can be seen to be localized near the bottom disk, with only small local strain rates for near the surface. For more details on the 3D shear band location, see [18] and references therein.

We now turn to the rheology of our immersed probes for $T \neq 0$. We first establish that for $\Omega = 0$, we recover the typical yield stress behavior of granular media. As indicated in fig. 2(b), experiments where we slowly ramp up $T$ then reveal a yield torque $T_y = 11.7 \pm 0.3 \text{ mNm}$, above which there is rapid flow, and below which there is, in essence, no flow.$^1$

To extract the static friction coefficient $\mu_s$ from the observed yield torque, we make the following three standard assumptions for frictional granular media: i) the pressure is hydrostatic: $P = \rho gz$, where for our material $\rho = 1850 \text{ kg/m}^3$; ii) the shear stresses are maximal, and the slip flow localizes at the cylindrical hull of the probe (area $2\pi r_c h_e$), where $r_c$ and $h_e$ are the effective dimensions of the probe—from earlier work, we know we can estimate $r_c$ and $h_e$ by extending the bare probe dimensions by half a grain diameter, i.e., $r_c = r + 0.5 \text{ mm}$ and $h_e = h + 1 \text{ mm}$ [$7$]; iii) the yielding shear stress is proportional to the product of friction coefficient and pressure: $\sigma_{\tau\phi}(z) = \mu_s P(z) = \mu \rho gz$. We find that

$$T = \int_A dA \sigma_{\tau\phi} = 2\pi r_c^2 \mu \rho g z_m h_e \Rightarrow$$

$$T/\lambda = \mu \rho g z_m = \mu P,$$

where $\lambda := 2\pi r_c^2 h_e$. (3)

Here $\lambda$ collects all geometric factors needed to relate torques to stresses (when the probe is not fully submerged, we need to adapt $h_e$). Substituting the relevant quantities for the experiments shown in fig. 2, $r_c = 20.5 \text{ mm}$, $h_e = 21 \text{ mm}$, and $z_m = 20 \text{ mm}$, we estimate that $\mu = 0.58 \pm 0.1$—consistently with independent measurements of $\mu$ for these beads which are in the range 0.52–0.58 [2,3].

Secondary rheology. — Now we turn our attention to the secondary rheology when both the imposed flow rate $\propto \Omega$ and the driving torque of the vane, $T$, are nonzero. Experimentally, we typically control $\Omega$ and $T$ and then measure the rotation rate $\omega$. We found that after a short transient, the rotation rate of the probe $\omega$ quickly reaches a steady state, and will show below that experiments where we control $\omega$ and measure $T$ yield identical results. We have also checked that runs done by sweeping the torque up or down are identical, and we have seen no history dependence, ruling out thixotropic or other more complex rheologies. Hence, $\omega(T)$ is a well-defined function fully determining the steady-state secondary rheology once the experimental parameters $\Omega$, $z_m$, $H$ and probe size are set.

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$^1$If we quench $T$ from zero to a value close to $T_y$, we can detect a very slow creep flow that has a rate that decreases in time and tends to zero.
In fig. 3 we show examples of such a rheological curve, \( \omega(T) \), for a 20 \times 20 mm\(^2\) probe buried at depth \( z_m = 20 \text{ mm} \), filling height \( H = 60 \text{ mm} \) and a range of rotation rates \( \Omega \). As expected \([6,7,15-17]\), the overall driving rate \( \Omega \) sets the only relevant timescale, so that \( \omega \propto \Omega \), and all rheological curves collapse when plotted as \( \omega/\Omega \) vs. \( T \).

Our data shows that there is a clear difference between coflow (\( \omega/\Omega > 0 \)) and counterflow (\( \omega/\Omega < 0 \)), and moreover suggests that for any finite value of \( \Omega \), the yield stress vanishes so that there is no extended range of torques where the probes rotation rate \( \omega = 0 \). To clarify these points, we present in fig. 4 a plot of \( \omega/\Omega \) for \( \Omega = 0.05 \text{ rps} \) on a log scale. Figure 4(a) confirms that indeed there is not a finite yielding threshold, as there is no extended “dead zone” where \( \omega = 0 \).

Note that the precession flow induced by the rotation of the driving disk (fig. 2(a)) means that for \( T = 0 \), the vane rotation rate is equal to the precession rate and, indeed, for our data shown in fig. 4(a), we find that \( \omega/\Omega \) is of order \( 10^{-3} \) for \( T = 0 \), consistent with fig. 2(a). To hold the vane completely fixed in the lab frame, we need to apply a negative “holding” torque \( T_H \approx -5.1 \text{ mN/m} \). In fig. 4(a) we also show that data obtained by torque control, i.e., setting \( T \) and measuring \( \omega \) (full symbols), and by rate-control, i.e., setting \( \omega \) and measuring \( T \) (open symbols), are identical within the experimental reproducibility.

Figure 4(a) reveals that in very good approximation, \( \omega \) varies as the sum of two exponential functions of \( T \), and all our data can be fitted extremely well by the following form:

\[
\omega/\Omega = C \left( e^{\frac{T-T_{\text{H}}}{T_+}} - e^{\frac{T-T_{\text{H}}}{T_-}} \right). \tag{4}
\]

Here \( T_H \) is the holding torque, and \( T_+ \) and \( T_- \) determine the exponential growth of \( \omega \) with torque for co- and counterflow. We note that the prefactor \( C \) is directly related to physical parameters \( \omega_p, T_H, T_+ \) and \( T_- \); by substituting \( T = 0 \), we immediately find

\[
C = \frac{\omega_p/\Omega}{e^{-T_H/T_+} - e^{-T_H/T_-}}. \tag{5}
\]

**Symmetry breaking.** — Our rheological data clearly shows that there is a significant difference between co- and counter-rotation of the probe and disk. This difference is easy to understand qualitatively, as the rotation of the bottom disk breaks the chiral symmetry. The 3D nature of our flow allows us to probe the rheology in cases when the induced flow by the probe is neither opposite nor parallel to the disk-induced anisotropy, by using a “perpendicular” orientation of the probe as indicated in fig. 1(c).

As shown in fig. 4(b), there is now no symmetry breaking between left- and right-handed rotation of the probe. In good approximation, the rotation is still exponential in \( T \), and a fit to eq. (4) yields that for this case \( T_H \approx 0 \text{ mN/m} \), whereas \( T_- \approx T_- \approx 0.65 \text{ mN/m} \).

We now return to the broken symmetry case. How to understand the difference between co- and counterflow? A subtlety in our experiment is that a priori it is not clear whether co- and counter-rotation should be defined with respect to the global disk rotation \( \Omega \) (i.e., \( \omega < 0 \) or \( \omega > 0 \)) or the local precession \( \omega_p \) (i.e., \( \omega < \omega_p \) or \( \omega > \omega_p \)). Our fitting form, eq. (4), tacitly assumes that co- and counter-rotation should be defined with respect to the disk rotation, \( \Omega \). Alternatively, we could first correct for the local precession \( \omega_p \), but as fig. 4(c) shows, defining co- and counterflow with respect to the precession rate \( \omega_p \), by plotting \( (\omega - \omega_p)/\Omega \), hides the simple torque dependence shown in fig. 4(a). A crucial aspect of fig. 4(a) is thus that it shows that when the torque is between \( T_H \) and 0, so that the vane rotates counter to the local precession but with the disk, the secondary rheological \( \omega(T) \) curve has the same log-slope as for \( T > 0 \) — the local precession rate is thus irrelevant for the resistance to flow. Only when the vane
starts to move against the global disk rotation for \( T < T_H \), this slope changes. Consistently with earlier work \[19\], moving with the fabric (i.e. co-rotation) is tougher than moving against the fabric (counter-rotation). We conclude that co- and counter-rotation should be defined with respect to the disk rotation, and not to the local precession rate.

To understand this, let us consider in more detail the flow in the “mixed” regime where \( T_H < T < 0 \), and the vane rotates against the local precession but the with the disk rotation. We know from experiments in 2D \[22\] and 3D \[19,20\] that persistent shear leads to the buildup of an anisotropic fabric. This anisotropy is encoded in both the contact and force structure in the material, and leads to a fragile structure that resists the current flow direction strongly, as revealed by experiments where reversing the flow direction leads to a transient in which the material has little resistance against flow, before restructuring of the material with the opposite anisotropy leads to a steady state of high resistance.

In our experiments, we separate the generation of the main flow and the rheological probing. Our data is consistent with a picture in which the rotation of the bottom disk creates an anisotropic fabric in the granular material, which ultimately leads to a difference in resistance to coflow and counterflow, and where the handedness of the anisotropy is set by the main flow, and thus determined by the sign of \( \Omega \), even if the vane moves against the local precession rate. So as long as \( \omega/\Omega > 0 \), the vane rotates against the anisotropy, even if \( \omega < \omega_p \). In other words, the residual precession rate is irrelevant for the rheology experienced by the probe, although it does, of course, set the vane rotation for \( T = 0 \). By separating driving and probing, we thus can set up a steady-state rotation of the probe with or against the anisotropic fabric, where the fabric is determined by the disk rotation only.

**Effect of depth.** — By changing the depth of the probe, \( z_m \), we can probe the effect of pressure on the rheology, and probe how the holding torque \( T_H \) and the characteristic torques \( T_a \) and \( T_c \) scale with pressure. Of course, changing the depth of the probe also changes the location with respect to the shear band, but we previously showed that the effective viscosity of a freely sinking probe only has a very weak dependence on \( z \) for \( H = 60 \text{ mm} \) —precisely the value we use here. Hence, changing \( z \) allows us to probe whether these characteristic torques have a frictional nature —in the sense that shear and normal stresses are proportional; if so, the torques, which can be translated into shear stresses (see eq. (3)), should be proportional to the pressure \( P \propto z_m \).

In fig. 5(a) we show flow profiles for a range of values of \( z_m \). The deeper the probe, the more to the left the curves shift, and the holding torque, where \( \omega \) goes through zero, shows a systematic variation with probe depth \( z_m \). In a frictional picture, one would expect that \( T_H \) would be proportional to the hydrostatic pressure \( \sim z_m \), and taking the same corrections for the finite size of the probe into account as in our calculation of the static friction coefficient, we would thus expect that \( T_H / \lambda = \mu_H P \), where \( \lambda \) is defined as in eq. (2), and \( \mu_H \) is an effective friction coefficient that determines the relation between holding torque and pressure.

As we shown in fig. 5(b), the holding torque is indeed approximately linear in \( z_m \), and also varies systematically with the probe size. When plotted as a function of \( P \), the data for \( T_H / \lambda \) for all our probes collapse on a single straight curve, with slope \( \mu_H \approx 0.28 \). The interpretation of this friction-like coefficient \( \mu_H \) is not straightforward. We first note that it is roughly a factor two smaller than the static coefficient. It is also different from an ordinary dynamical friction coefficient, which would measure the resistance to flow of a granular medium, whereas here
we measure the resistance to secondary flow in an already flowing granular medium. Notwithstanding the precise value of \( \mu \), what is clear is that the holding torques are proportional to the pressure — the rheology of secondary granular flows has a frictional character.

Motivated by this success, we plot in fig. 5(c) the values of \( T_\perp/\lambda \) and \( T_\parallel/\lambda \) as a function of \( P \). Interestingly, we find that \( T_\parallel \) is also proportional to \( P \) as \( T_\parallel/\lambda \approx 0.03P \), whereas \( T_\perp \) goes to a finite intercept for \( P \to 0 \), and \( T_\perp/\lambda \approx 0.02 + 0.03P \). This implies that there is a significant qualitative difference between co- and counterflow.

Discussion and outlook. — Our study shows that for secondary rheology, the anisotropy is a key parameter. Surprisingly, the difference between co- and counterflow is not only quantitative, but qualitative: In counterflow, the characteristic stresses \( T_\parallel \) and \( T_\perp \) are both proportional to the pressure, whereas for coflow, the characteristic torque \( T_\parallel \) is not simply proportional to \( P \). This surprising finding, and the difference between the scaling of \( T_\parallel \), \( T_\perp \) and \( T_\perp \) beg for further detailed (numerical) studies of 3D secondary granular rheologies.

Notwithstanding the complex variation of our characteristic torques with pressure, we note that our secondary rheology can be written as \( \omega(z)/\omega(z) = F(T/P, P) \) — for counterflow, we get in good approximation that \( \omega(z)/\omega(z) = F(T/P) \), so that the local flow is encoded in \( \omega(z) \), whereas the secondary rheology is frictional. Our scaling for \( T_\perp \) and \( T_\parallel \) suggests that for deeper layers, where the finite intercept of \( T_\parallel \) becomes of less relative importance, we approach \( T_\perp \approx T_\parallel \sim P \) — so the effect of symmetry breaking diminishes with distance to the source of the main flow.

We note here that for the “perpendicular” probing, i.e., a measurement protocol not sensitive to the flow direction, the holding torque is zero, and so in these cases it is more difficult to observe the frictional aspects of the secondary rheology.

Our secondary rheology has a rate which is clearly exponential in the shear stresses — consistently with [6] and [10], but differently from [7]. We have no good explanation for this, but point out that the linear regime of [7] was only observed for slow secondary flows, and it is conceivable, that the observed linear regime simply corresponds to the overlap of two exponential functions.

We note that in [6], the yield stress was used as characteristic scale, while here we use the “holding” stress. We expect that the results of [6] could be recast using the holding stress, as for counterflow, all relevant stress scales are proportional to the pressure, and so are the yield and holding stresses. For the case of coflow, not probed in [6], we found here the situation to be more complex. We have no clear picture of what causes the stress scale given by the finite intercept of \( T_\parallel \) in fig. 5(c) for coflow, but note here that differences between the secondary rheology for co- and counterflow have also recently been observed in numerical studies [23].

Open questions for the future are to disentangle the role of pressure and distance to the source of the main flow, and models for the role of anisotropy in slow granular flows.

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We acknowledge technical assistance of JEROEN MESMAN. EW acknowledges funding by the Dutch funding agency FOM.

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